

SHORT CONTRIBUTION

The dependence of the power-law exponent on surface roughness and stability in a neutrally and stably stratified surface boundary layer

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RESUMEN

La dependencia del exponente (p) de la ley exponencial del perfil del viento de la rugosidad superficial y de la estabilidad atmosférica se describe suponiendo una variación con la altura de ϕ (Huang, 1979; Hanafusa *et al.*, 1986) y empleando la formulación de Zilitinkevich (1989) para establecer los perfiles de velocidad en una capa límite planetaria neutra y establemente estratificada.

Las estimaciones teóricas del exponente de la ley exponencial son perfectamente comparables con los datos de dicha ley provenientes de varias fuentes y de los análisis teóricos de Irwin (1979). Se reconoce que se necesita más trabajo antes de que la validez de la metodología propuesta pueda ser plenamente establecida.

ABSTRACT

The dependence of the wind profile power-law exponent (p) on surface roughness and atmospheric stability is depicted assuming a height variation of p (Huang, 1979; Hanafusa *et al.*, 1986) and using the formulation of Zilitinkevich (1989) for determining the velocity profiles in a neutrally and stably stratified planetary boundary layer. The theoretical estimates of the power-law exponent compare perfectly with the power-law exponent data from various sources and the theoretical analysis from Irwin (1979). It is recognized that more work is necessary before the validity of the suggested methodology can be fully established.

1. Introduction

The magnitude and direction of near-surface winds and their variations with height, in the planetary boundary layer, are of considerable interest in both research and applied studies, such as Air Pollution (Stern, 1976; Hanna *et al.*, 1982), Building Design (Sachs, 1978; Melaragno, 1982) and Wind Energy Technology (Justus *et al.*, 1976; Gourieres, 1982). Since most of the available wind speed measurements have been made close to the ground, it is necessary to extrapolate the wind speed profiles within the surface boundary layer (Segal and Pielke, 1988). The most common extrapolation is based on the power-law velocity equation, which is preferred by engineers, for mathematical simplicity, and also provided a reasonable fit to the observed wind velocity profiles, in the lower part of the planetary boundary layer. It may be expressed as,

$$u(z)/u_1 = (z/z_1)^p, \quad p > 0 \quad (1)$$

where $u(z)$ and u_1 are the wind speed at the heights z and z_1 , respectively, where the height

z_1 is at the reference level. The exponent p is then to be determined on the basis of surface configuration and atmospheric conditions.

As discussed by Arya (1988) the logarithmic profile law has a sounder theoretical and physical basis than the power-law profile, especially within the neutral surface layer. There is no exact correspondence between the power-law and logarithmic wind profile, because the two profile shapes are different. However, the use of the power law gives a reasonably accurate representation of the wind speed profiles and a better representation than does the logarithmic law for the entire depth of the planetary boundary layer (Counihan, 1975; Hanafusa *et al.*, 1986; Segal and Pielke, 1988). On the other hand, the power-law exponent, p , is found to depend on the surface roughness, atmospheric stability and height range over which the power-law is fitted to the observed profile. Following Panofsky and Prasad (1965) and Panofsky and Dutton (1983), Equation (1) can be rearranged to give,

$$p = (z/u)(\partial u/\partial z) \quad (2)$$

Given the dimensionless wind shear which is a function of $\zeta = z/L$ only,

$$\partial u/\partial z = (u_*/kz)\phi(z/L) \quad (3)$$

integration of Equation (3) gives the surface layer wind profile for the stable atmospheric conditions (Businger *et al.*, 1971; Zoumakis and Kelessis, 1991a),

$$u(z) = (u_*/k)\{\ln(z/z_o) + \beta[(z - z_o)/L]\} \quad (4)$$

where u_* is the friction velocity, $k = 0.40$ is the von Karman constant, L is the Monin-Obukhov stability length, z_o is the surface roughness length (which is a surface characteristic of a fixed site and is not supposed to change with stability), and $\beta = 4.7$. Substitution of Equations (3) and (4) into (2) finally yields an approximate relationship:

$$p = [1 + 4.7(z/L)]/[\ln(z/z_o) + 4.7(z/L)] \quad (5)$$

The value of p is computed at a height z ; however it provides, as suggested by Huang (1979) and Sedefian (1980), a good representation for a sublayer between z_1 and z_2 when $z = (z_1 z_2)^{0.5}$. Unfortunately, the theoretical predictions of the power-law exponent from Equation (5), under stable conditions, overestimate the p -values (Irwin, 1979; Zoumakis and Kelessis, 1991b; Zoumakis, 1992).

2. Methodology and Discussion

For a comparison of the power-law and surface layer wind profiles, it was assumed that the near-surface velocity profile should satisfy the logarithmic + linear law (4); however, the region where it holds true is limited in addition to the inequality $z \ll h$, also by another inequality $L \ll h$ (Zilitinkevich, 1989), where h is the stably stratified planetary boundary layer (*PBL*) depth (Delage, 1974; Wyngaard, 1975; Arya, 1977; Brost and Wyngaard, 1978; Yamada, 1979; Zilitinkevich, 1989). On the other hand, in the evening and at night with the increase of hydrostatic stability the atmospheric *PBL* collapses; its upper boundary falls more and more, while turbulence in the overlying region degenerates (Zilitinkevich, 1989). A further analysis involved substituting the surface layer wind profile from Equation (4) by a velocity profile for the whole

PBL region, $z \leq h$,

$$u(z) = (u_*/k)[\ln(z/z_0) + f_u(z/h, \mu)], \quad v(z) = (-u_*/k)f_v(z/h, \mu)\text{sign}(f) \quad (6)$$

$$L = -u_*^3/kB_s, \quad \mu = ku_*/|f|L \quad (7)$$

$$h = u_*\Lambda(\mu)/|f| \quad (8)$$

where μ is the dimensionless stratification parameter, B_s is the near-surface value of the vertical turbulent buoyancy flux, f is the Coriolis parameter and the functional forms for $f_u(z/h, \mu)$, $f_v(z/h, \mu)$ and $\Lambda(\mu)$ are taken from Zilitinkevich (1989). On the other hand, the velocity profiles from Equations (6), (7) and (8) satisfy the logarithmic + linear law (4) and $v(z) = 0$, in the region at $z \ll h$ and $L \ll h$. Combining Equations (1), (6), (7) and (8), the wind profile power-law exponent p , for the whole *PBL* region $z \leq h$, has the following form:

$$p(z_1, z_2, z_0, L, h) = \ln(U_2/U_1)/\ln(z_2/z_1) \quad (9)$$

where U_2 and U_1 are the wind vectors at the heights z_2 and z_1 , respectively (Ku *et al.*, 1987). Assuming an independent height scale h (Clarke and Hess, 1974; Zilitinkevich and Deardorff, 1974), Figure 1 compares the theoretical predictions for p , from Equation (9), with the power-

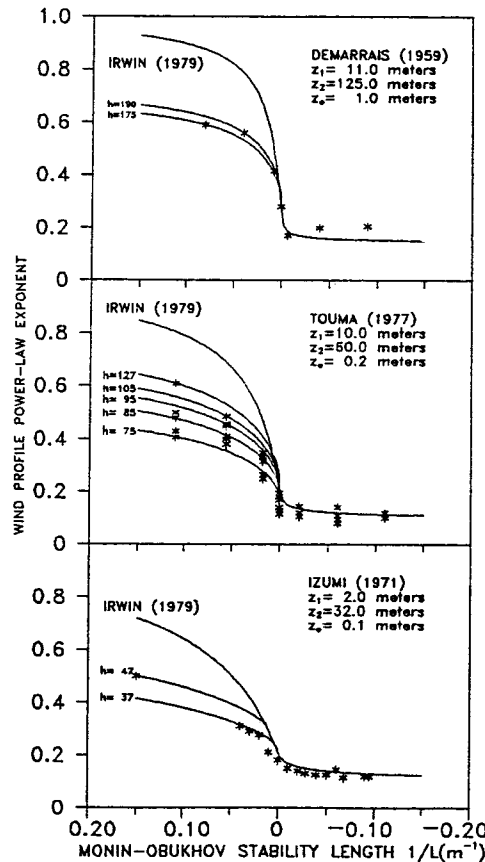


Fig. 1. Comparisons of the theoretical estimates of the power-law exponent values, from Equation (9), with experimental data from three sources and theoretical predictions from Irwin's formula.

law exponent data (De Marrais, 1959; Izumi, 1971; Touma, 1977; Irwin, 1979) and estimates from Equation (5). In Figure 1 the plots indicate that, during stable conditions, the theoretical estimates from Equation (9) compare perfectly with the data. Despite the extremely low h -values used in Figure 1 and the large uncertainty in determining the height scale h , the suggested Equation (9) seems to explain the dependence of p on surface roughness and atmospheric stability, during stable conditions.

Although the mechanism is not completely clear at this stage, in view of Equation (9) and the power-law exponent data published by Irwin (1979), it appears possible that the classical Equation (5) overestimates the p -values. However, more theoretical studies and more data on the power-law exponent as a function of L and h , at different Pasquill stability classes, over representative land types, will be required before the suggested procedure can be used in an operational mode. Hopefully, future work will help us to increase our understanding of the dependence of the power-law exponent on surface roughness and atmospheric stability, during stable conditions.

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